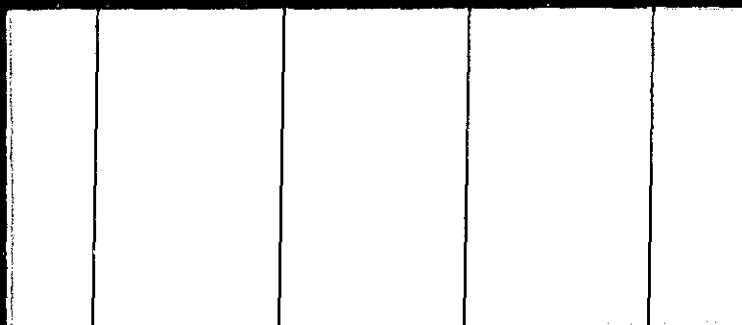


General Disclaimer

One or more of the Following Statements may affect this Document

- This document has been reproduced from the best copy furnished by the organizational source. It is being released in the interest of making available as much information as possible.
- This document may contain data, which exceeds the sheet parameters. It was furnished in this condition by the organizational source and is the best copy available.
- This document may contain tone-on-tone or color graphs, charts and/or pictures, which have been reproduced in black and white.
- This document is paginated as submitted by the original source.
- Portions of this document are not fully legible due to the historical nature of some of the material. However, it is the best reproduction available from the original submission.



(NASA-CR-142759) THE INVESTIGATION OF O AND
N2 DENSITIES FROM THE OSO-7 EXTREME UV DATA
Final Report (Wolf Research and Development
Corp.) 36 p HC \$3.75

CSCL 03B

N75-24644

G3/92

Unclas
17689

THE INVESTIGATION OF O AND N₂
DENSITIES FROM THE OSO-7
EXTREME UV DATA
FINAL REPORT

Submitted To: NASA
Code PY
Headquarters
Washington, D.C.

Submitted By: Wolf Research and Development Group
6801 Kenilworth Avenue
Riverdale, Maryland 20840

Date Submitted: February 3, 1975

Reference: WOLF No. 00079

TABLE OF CONTENTS

	<u>Page</u>
1.0 INTRODUCTION	1
2.0 ANALYSIS DESCRIPTION	2
3.0 ORBITAL AND GEOMETRICAL ANALYSIS	10
4.0 RESULTS OF THE ANALYSIS	17
5.0 DISCUSSIONS	29
REFERENCES	32

SUMMARY

This report presents the results of a study of solar radiation observations in the extreme ultra violet spectrum from 200 to 600 Å made by the OSO-7 satellite. The results of special interest in this work were those made under the influence of attenuation by the atmosphere in the 250 to 500 km. altitude range.

Using published molecular absorption cross-sections at 304 Å and 256 Å, we were able to validate the Jacchia atmospheric model (1971) and showed that a mean exospheric of 1050°K was appropriate for the sunset data. The cross-sections obtained by Hinteregger and Hall (1969) were shown to be most appropriate at 256 Å, whereas those of Knight, Uribe and Woodgate (1973) were employed for 304 Å. Employing the validated atmospheric model we were then able to derive values of 6.3 ± 0.6 Barns and 6.9 ± 3.0 ~~Barns~~ for the cross-sections of O and N₂, respectively at 285 Å and 7.6 ± 0.8 ~~Barns~~ and 4.8 ± 3.9 ~~Barns~~ for the respective cross-sections at 355 Å.

SECTION 1.0

INTRODUCTION

Wolf Research and Development Corporation is pleased to submit this final report on project NAS W-2595 to NASA Headquarters. The report describes the final results obtained in analyzing OSO-7 data to evaluate atmospheric absorption cross-sections in 300-500 km. altitude range and the 200-400 Å wavelength range.

All the analysis and results pertinent to the conclusions are presented in this report. In Section 2.0 the basic intensity analysis is given, followed by the geometrical and orbital analysis in Section 3.0. Section 4.0 gives a discussion of the results analysis and conclusions are presented in Section 5.0.

SECTION 2.0

ANALYSIS DESCRIPTION

The OSO-7 experimental tapes contain information on the date of acquisition of solar intensity data, orbital characteristic information and the observed intensities at one or two lines in the spectrum. The experimental data files are organized into pages or frames, there being 96 intensities collected over 15.36 seconds for each line being studied presented within each frame. The fundamental sampling interval was 160 milliseconds and the data appeared as photon counts.

For the purposes of our analysis the section of the quantity of data to average before applying an inversion analysis is indicated by the requirement for a suitable compromise between minimizing random errors and systematic errors. The greater the averaging time, the smaller the random error, but the larger the systematic error introduced by non-linearities in the model. In the work reported here we assumed that the average intensity at the midtime of a 5 second data span is sufficient. Thus we have averaged the intensity for three sections, each 32 points long, within each of the frames and printed out as shown in Figure 2.1. In the Figure 2.1, the various columns are as follows:

Frame: Experimental Frame Number

Midtime: Middle time of the experimental frame (seconds from start of day, Universal Time).

AVEUV1: Average intensity in the medium wavelength of the frame.

AVEUV2: Average intensity in the long wavelength of the frame.

ORIGINAL PAGE IS
OF POOR QUALITY

FRAME	MIDTIME	AVEU01	AVEU02	X	Y	Z	ALTITUDE	SUN ANGLE
1	31056.659	2.198	2.683	-0.3145028146J	04 -0.4757332637D	04 -0.3640466719D	04 0.4012653938D	03 -0.2406122379D
2	31072.219	2.845	3.302	-0.3043358715D	04 -0.4323324364D	04 -0.3652461600D	04 0.4039976077D	03 -0.2238773939D
3	31087.573	7.760	6.208	-0.2940764553D	04 -0.4388434435D	04 -0.3663346006D	04 0.4049259464D	03 -0.2070334815D
4	31102.938	17.354	10.704	-0.2837277077J	04 -0.4743395553D	04 -0.3673116585D	04 0.4057481154D	03 -0.1922357354D
5	31119.297	27.792	16.729	-0.2738608703J	04 -0.4972235815D	04 -0.3681366155D	04 0.4086317792D	03 -0.1742491889D
6	31133.650	37.313	21.385	-0.2633494083J	04 -0.5004314160D	04 -0.3688177367D	04 0.4104904217D	03 -0.1573128576D
7	31149.016	45.917	25.406	-0.2527581615D	04 -0.5095947274D	04 -0.3695469983D	04 0.4123216399D	03 -0.1403381577D
8	31164.379	54.771	27.583	-0.2420876472J	04 -0.5147319465D	04 -0.3700843226D	04 0.4141539880D	03 -0.1239259844D
9	31179.738	56.667	30.417	-0.2313516699J	04 -0.5193746933D	04 -0.3705119929D	04 0.4161916445D	03 -0.1062858362D
10	31195.078	63.073	33.167	-0.2205426753J	04 -0.5242715010D	04 -0.3708287255D	04 0.4180524281D	03 -0.08922824608D
11	31210.457	67.510	34.458	-0.2096670883J	04 -0.5283100104D	04 -0.3710313682D	04 0.4199030537D	03 -0.0721319360D
12	31225.816	69.535	35.698	-0.1987282262J	04 -0.5331335340D	04 -0.3711218574D	04 0.4217412106D	03 -0.05506595408D
13	31241.176	69.302	36.021	-0.1877331616J	04 -0.5374302173D	04 -0.3711952440D	04 0.4237952291D	03 -0.03796670074D
14	31256.539	67.474	36.833	-0.1756767194J	04 -0.5414110565D	04 -0.3709735724D	04 0.4256997359D	03 -0.0206636440D
15	31271.878	65.104	36.127	-0.1655992634J	04 -0.5453176659D	04 -0.3707331476D	04 0.4275119109D	03 -0.003774329816D
16	31287.258	63.479	35.781	-0.1546133460J	04 -0.5497119110D	04 -0.3703750272D	04 0.4293472224D	03 0.01350922995D
17	31302.617	61.250	34.894	-0.1432124760J	04 -0.5537122445D	04 -0.3699136066D	04 0.4314063760D	03 0.03039753585D
18	31317.977	58.375	33.573	-0.1319679143J	04 -0.5580117735D	04 -0.3693377540D	04 0.4332655103D	03 0.04745270625D
19	31333.336	57.240	31.646	-0.1208837929J	04 -0.5593613090D	04 -0.3686509861D	04 0.4351090394D	03 0.06446219302D
20	31348.699	55.969	32.198	-0.1093606550J	04 -0.5524272944D	04 -0.3678532927D	04 0.4369354317D	03 0.08148728131D
21	31364.059	54.855	30.000	-0.9765566053J	04 -0.5653254021D	04 -0.3669177921D	04 0.4388891674D	03 0.09896551260D
22	31379.418	52.729	29.146	-0.8627290841J	04 -0.5681534271D	04 -0.3658170016D	04 0.4405918537D	03 0.1158910403D
23	31394.777	50.969	29.573	-0.7486445249J	04 -0.5707109123D	04 -0.3647664317D	04 0.4424719642D	03 0.1327726303D
24	31410.137	49.406	28.500	-0.6373007329J	04 -0.5733555570D	04 -0.3635750346D	04 0.4444709380D	03 0.1490760941D
25	31425.496	48.956	28.354	-0.5234162657J	04 -0.5752901581D	04 -0.3622316536D	04 0.4462564953D	03 0.1658555256D
26	31440.859	49.698	27.354	-0.4087480539J	04 -0.5773329785D	04 -0.3607795054D	04 0.4480181617D	03 0.1825803020D
27	31456.219	49.125	26.427	-0.2939894034J	04 -0.5792124014D	04 -0.3592197367D	04 0.4497527272D	03 0.19992385753D
28	31471.574	49.323	26.844	-0.1791386517J	04 -0.5803542550D	04 -0.3578508440D	04 0.4519211370D	03 0.2158318549D
29	31486.938	47.958	25.490	-0.6424433311D	04 -0.5825042454D	04 -0.3557949578D	04 0.4536785925D	03 0.2323479355D
30	31502.297	48.656	25.490	-0.5066637603J	04 -0.5833577307D	04 -0.3539179439D	04 0.4554096665D	03 0.2487864629D
31	31517.656	48.948	24.938	-0.1655913623J	04 -0.5850311585D	04 -0.3519353624D	04 0.4571121145D	03 0.2651427140D
32	31533.016	47.598	23.938	-0.2904434717J	04 -0.5861247666D	04 -0.3498607931D	04 0.4592530828D	03 0.2814162496D
33	31548.379	47.135	24.063	-0.3555211708J	04 -0.5886744790D	04 -0.3476716543D	04 0.4609756650D	03 0.2975974429D
34	31563.735	46.625	23.844	-0.5059709111J	04 -0.5876493763D	04 -0.3453770710D	04 0.4625696071D	03 0.3136784151D
35	31579.098	47.771	22.931	-0.6243557621D	04 -0.5881494946D	04 -0.3429850983D	04 0.4643331312D	03 0.3296587594D
36	31594.457	46.063	21.281	-0.7387326733J	04 -0.5885215164D	04 -0.3405080970D	04 0.4664368209D	03 0.3455387013D
37	31609.816	47.333	21.823	-0.8524817459J	04 -0.5885792719D	04 -0.3379062431D	04 0.4681168246D	03 0.3613036827D
38	31625.176	46.323	20.740	-0.9664721673J	04 -0.5896325333D	04 -0.3352115686D	04 0.4697667896D	03 0.3769547087D
39	31640.539	45.927	21.094	-0.1080213660J	04 -0.5894712285D	04 -0.3324168731D	04 0.4713255565D	03 0.3924913604D
40	31655.899	44.719	19.813	-0.1192550714J	04 -0.5891518769D	04 -0.3295357089D	04 0.4734424516D	03 0.4079066866D
41	31671.258	44.365	20.563	-0.1306422294J	04 -0.5893751364D	04 -0.3265486627D	04 0.4750733286D	03 0.4231827296D
42	31686.617	44.760	19.208	-0.1418917618J	04 -0.5896713256D	04 -0.3234652007D	04 0.4766739227D	03 0.4383407977D
43	31701.977	44.313	20.104	-0.1530984800J	04 -0.5896337556D	04 -0.3202862474D	04 0.4782406262D	03 0.4533570840D
44	31717.336	43.104	19.167	-0.1642749120J	04 -0.5895264282D	04 -0.3170228375D	04 0.4802415336D	03 0.4652397815D
45	31732.699	44.802	18.635	-0.1753923867J	04 -0.5893880196D	04 -0.3136577086D	04 0.4818185712D	03 0.4829765318D
46	31748.059	44.093	18.719	-0.1864551499J	04 -0.5892418004D	04 -0.3102011243D	04 0.4833626673D	03 0.4975621199D
47	31763.418	42.927	18.385	-0.1974627097J	04 -0.5890836472D	04 -0.3066332749D	04 0.4843721877D	03 0.5119963564D
48	31778.777	43.740	17.927	-0.2094303523J	04 -0.5897157653D	04 -0.3030238773D	04 0.4860284625D	03 0.5262815456D
49	31794.137	43.948	17.688	-0.2193206184J	04 -0.5877271741D	04 -0.2992924930D	04 0.4883246009D	03 0.5404010106D
50	31809.496	43.655	18.333	-0.2301462750J	04 -0.5875206399D	04 -0.2954868393D	04 0.4899071394D	03 0.5543576037D
51	31824.859	43.646	18.365	-0.2409060693J	04 -0.5872716662D	04 -0.2915867825D	04 0.4912544765D	03 0.5681510727D
52	31840.219	44.156	17.740	-0.2516150850J	04 -0.5870610931D	04 -0.2875992722D	04 0.4931176514D	03 0.5817772636D
53	31855.578	43.917	18.010	-0.2622211017J	04 -0.5868347929D	04 -0.2833542696D	04 0.4945678868D	03 0.5952226664D
54	31870.938	43.948	18.073	-0.2727699855J	04 -0.5865318778D	04 -0.2794932196D	04 0.4959836442D	03 0.6084906101D
55	31886.297	44.104	17.250	-0.2832280384J	04 -0.5862321255D	04 -0.2751620946D	04 0.4973630394D	03 0.6215775834D
56	31901.656	44.271	17.832	-0.2936286111J	04 -0.5859341379D	04 -0.2704558764D	04 0.4991452021D	03 0.6344869253D
57	31917.016	44.948	17.917	-0.3034218150J	04 -0.5861379617D	04 -0.2664672155D	04 0.5005244951D	03 0.6472014190D
58	31932.379	44.354	18.563	-0.3141287102J	04 -0.5852334610D	04 -0.2614997610D	04 0.5018299516D	03 0.6597279531D
59	31947.738	44.042	19.438	-0.3242416622J	04 -0.5843263967D	04 -0.2574571133D	04 0.5031760457D	03 0.6720568809D
60	31963.098	43.875	18.906	-0.3342847538J	04 -0.5840303972D	04 -0.2528495360D	04 0.5043604317D	03 0.6841951179D
61	31979.457	43.638	18.094	-0.3442095120J	04 -0.5841732153D	04 -0.2491533871D	04 0.5056173450D	03 0.6961255089D

FIGURE 2.1. COMPUTER OUTPUT SHOWN AVERAGED INTENSITY AND ORBITAL DATA

$\left. \begin{matrix} X \\ Y \\ Z \end{matrix} \right\} :$
Satellite geocentric coordinates in km. at the midtime of the experimental frame.

Altitude: Altitude of the satellite from the Earth's surface in km. at the midtime of the experimental frame.

Sun Angle: Cosine of the angle between the lines joining the satellite and solar center to the geocenter.

The calculation of the position, altitude and the Sun angle of the satellite will be discussed in detail in Section 3.0.

The intensity data at a given wavelength λ is measured by the satellite instrument. We assumed that when the path from the satellite to the Sun did not pass through any layer below 480 kms. there was no significant attenuation. Thus this data was accumulated and averaged to give the mean intensity per observation for zero attenuation, $I_{0\lambda}$.

When the path from the satellite to the Sun fell below 480 kms. the mean intensity per observation I_{λ} generally fell below $I_{0\lambda}$ indicating attenuation. The total optical thickness along the path through the atmosphere could be estimated as:

$$\tau_{\lambda} = \log_e \left(\frac{I_{0\lambda}}{I_{\lambda}} \right) \quad (2.1)$$

When the Sun-satellite line fell below the 300 km. altitude it was found that no significant signal was received. Thus the intensity yields information about attenuation in the 260 to 480 km. altitude range.

We divided the atmosphere in this altitude range into 22 layers each 10 kms. thick. This was considered to be the greatest degree of detail we could reasonably consider when using one file of data (one orbit) only.

The equation used for analyzing the data was

$$\tau_{\lambda j} = \sum_i F_{ij} \sigma_{i\lambda} \quad (2.2)$$

where

$\tau_{\lambda j}$ is the j^{th} observation of optical thickness at wavelength λ .

F_{ij} is the path length in kms. through the i^{th} layer for the j^{th} observation (see next section).

$\sigma_{i\lambda}$ is the mean absorption cross-section in the i^{th} layer (km^{-1}).

$\sigma_{i\lambda}$ is made up of two parts, one for absorption by O and the other for absorption by N_2 . If (n_{0i}, n_{Ni}) are the mean densities of O and N_2 , respectively in the i^{th} layer and $\sigma_{0\lambda}$ and $\sigma_{N\lambda}$ are the respective cross-sections at wavelength λ we have that

$$\sigma_{i\lambda} = n_{0i} \sigma_{0\lambda} + n_{Ni} \sigma_{N\lambda} \quad (2.3)$$

Thus the values of λ_i contain information about both the O and N₂ densities, as well as their cross-sections. Clearly, we need to know at least 3 of the 4 unknowns in each layer before the other parameter can be determined. Therefore, even if the cross-sections are known it is impossible to unambiguously derive both the N₂ and O densities independently.

The procedure we employed in our analysis was to use the standard atmospheres of Jacchia (1971) to determine the appropriate range of exospheric temperatures where the cross-sections were known (256 Å and 304 Å) and then to employ a least squares analysis to derive the cross-sections σ_{N_2} and σ_{O} at the wavelengths where they were not known (285 Å and 355 Å). The details of the results will be considered in Section 4.0 of this report.

In deriving the values of $\sigma_{i\lambda}$ using equations of the form of (2.2), it was necessary to determine the appropriate weighting factor to be used with each equation. The objective was to evaluate the variance, σ^2 , in τ_λ and then divide the corresponding equation by σ .

Suppose that $I_{0\lambda}$ was derived from a total of $N_{0\lambda}$ counts over a time $T_{0\lambda}$ and I_λ was derived from a total of N_λ counts over a time T_λ . Then

$$\left. \begin{aligned} I_{0\lambda} &= N_{0\lambda} T_{0\lambda}^{-1} \\ I_\lambda &= N_\lambda T_\lambda^{-1} \end{aligned} \right\} \quad (2.4)$$

Then

$$\left. \begin{aligned} V(I_{0\lambda}) &= N_{0\lambda} T_{0\lambda}^{-2} \\ V(I_{\lambda}) &= N_{\lambda} T_{\lambda}^{-2} \end{aligned} \right\} \quad (2.5)$$

give the variances in $I_{0\lambda}$ and I_{λ} , respectively, since $N_{0\lambda}$ and N_{λ} are Poisson counts.

It follows that

$$\left. \begin{aligned} I_{0\lambda}^{-2} V(I_{0\lambda}) &= N_{0\lambda}^{-1} \\ I_{\lambda}^{-2} V(I_{\lambda}) &= N_{\lambda}^{-1} \end{aligned} \right\} \quad (2.6)$$

and

$$V\left(\frac{I_{0\lambda}}{I_{\lambda}}\right) = \left(\frac{I_{0\lambda}}{I_{\lambda}}\right)^2 \left(\frac{1}{N_{0\lambda}} + \frac{1}{N_{\lambda}}\right) \quad (2.7)$$

Using (2.1) we then obtain the simple result,

$$V(\tau_{\lambda}) = \left(\frac{1}{N_{0\lambda}} + \frac{1}{N_{\lambda}}\right) \quad (2.8)$$

which is the required variance, σ^2 .

A series of experiments we performed on the inversion of overdetermined linear equations indicated that the best combination of accuracy and computer time requirement would be yielded by the IBM systems subroutines package inversion programs LLSQ and DLLSQ. In practical applications, however, these solution procedures failed to successfully invert (2.1) consistently. The reason for this failure is not clear but we believe it relates to the substantial random errors occurring in the τ_j estimates. The program was carefully checked to eliminate any possible dimensioning or other errors and the fact that the routine successfully inverted some data sets, supports the hypothesis that the routines were being employed correctly.

The lack of consistency in producing successful inversions seriously slowed our production analysis activity. As a consequence we developed a new procedure which has proved to be very stable and reliable, while producing sensible answers consistent with those generated by successful LLSQ runs. The procedure is based on the summation of sets of equations of the form (2.1) to generate a right upper triangular matrix on the left-hand side. This proved to be particularly easy to accomplish since the altitude of the grazing incidence increased monotonically during sunrise and decreased monotonically during sunset. (This altitude establishes the number of leading zero F_{ij} values in (2.1).) Thus the procedure required merely the summing of sets of adjacent model equations. Non-singularity was assured in all practical cases by the stipulation that there should be at least one grazing incidence in each layer of the atmospheric model. This is discussed more fully in Section 3.1.

Once the left-hand side matrix is right upper triangular, the cross-section in the top layer can be solved for immediately and successive back substitution employed

to sections within all other layers. In practice, we have found that the solution can be generated in much less time than by LLSQ since the procedure requires N^2 multiplications compared to $N^2(N+n)$ for the standard least squares procedure. (N is the number of coefficients and n the number of equations.) The results generated by the application of the new routine to the data analyzed in the previous report gave cross-section profiles very similar to those yielded in LLSQ.

SECTION 3.0

ORBITAL INTERPOLATION AND GEOMETRIC ANALYSIS

3.1 ORBITAL INTERPOLATION

The data tapes contain orbital data in geocentric coordinates at one minute intervals as illustrated in Figure 3.1 of the first quarterly progress report on this project. The problem that we considered was that of obtaining the geocentric coordinates at the mid points of the experimental frames using the given data at one minute intervals.

The first step was to compute the nearest given orbital data point to the frame mid point. Clearly this could not be more than 30 seconds away. We thus knew the orbital elements including velocity components at an epoch t and wished to compute the elements at a time $t + \Delta t$ where Δt was in the range -30 seconds to +30 seconds. Let $X_i(t)$ be the i th coordinate at time t . Cowell's equations for two body motion is

$$\ddot{X}_i = - \frac{X_i}{R^3} \quad (3.1)$$

where μ is the gravitational constant = $3.986013 \times 10^5 \text{ km}^3 \text{ sec}^{-2}$

and R is the geocentric radius

$$= (\sum_{i=1}^3 X_i^2)^{1/2}$$

In making the analysis we neglected the time variation of R since the OSO-7 orbit is not far from circular (eccentricity around 0.0165). Thus we obtained from equation (3.1)

ORIGINAL PAGE IS
OF POOR QUALITY

11

TAPE NUMBER= CA0614 FILE NUMBER= 62
FRAME NUMBER= 103 YEAR= 1972 DAY OF YEAR= 06 TIME OF DAY= 36180 SECON DS
MEDIUM WAVELENGTH= 314.90 ANGSTROMS LONG WAVELENGTH= 368.10 ANGSTROMS
X Y Z
SATELLITE POSITION -0.5645846567D 04 -0.2383256683D 04 -0.2786425880D 04 (KMS.)
SATELLITE VELOCITY -0.3846352428D 01 -0.3119421843D 01 -0.2795339348D 01 (KMS./SEC.)
SOLAR POSITION 0.7825655318D 00 -0.2211511357D 00 -0.5590211883D 01 (A.U.)
TAPE NUMBER= CA0614 FILE NUMBER= 62
FRAME NUMBER= 104 YEAR= 1972 DAY OF YEAR= 06 TIME OF DAY= 36240 SECON DS
MEDIUM WAVELENGTH= 314.90 ANGSTROMS LONG WAVELENGTH= 368.10 ANGSTROMS
X Y Z
SATELLITE POSITION -0.5401704529D 04 -0.2725814044D 04 -0.2947451874D 04 (KMS.)
SATELLITE VELOCITY 0.4280499935D 01 -0.3591240343D 01 -0.2569350094D 01 (KMS./SEC.)
SOLAR POSITION 0.9625689387D 00 -0.2211405183D 00 -0.5589749109D 01 (A.U.)
TAPE NUMBER= CA0614 FILE NUMBER= 62
FRAME NUMBER= 105 YEAR= 1972 DAY OF YEAR= 06 TIME OF DAY= 36300 SECON DS
MEDIUM WAVELENGTH= 314.90 ANGSTROMS LONG WAVELENGTH= 368.10 ANGSTROMS
X Y Z
SATELLITE POSITION -0.5132335754D 04 -0.3074142517D 04 -0.3094562836D 04 (KMS.)
SATELLITE VELOCITY 0.4693454023D 01 -0.3592049921D 01 -0.2331804663D 01 (KMS./SEC.)
SOLAR POSITION 0.9625720322D 00 -0.2211298477D 00 -0.9589286429D 01 (A.U.)
TAPE NUMBER= CA0614 FILE NUMBER= 62
FRAME NUMBER= 106 YEAR= 1972 DAY OF YEAR= 06 TIME OF DAY= 36360 SECON DS
MEDIUM WAVELENGTH= 314.90 ANGSTROMS LONG WAVELENGTH= 368.10 ANGSTROMS
X Y Z
SATELLITE POSITION -0.48386556ED 04 -0.3438233281D 04 -0.3227102448D 04 (KMS.)
SATELLITE VELOCITY 0.5083323419D 01 -0.3433601375D 01 -0.2083915591D 01 (KMS./SEC.)
SOLAR POSITION 0.9625751376D 00 -0.2211191733D 00 -0.9588877747D 01 (A.U.)
TAPE NUMBER= CA0614 FILE NUMBER= 62
FRAME NUMBER= 107 YEAR= 1972 DAY OF YEAR= 06 TIME OF DAY= 36420 SECON DS
MEDIUM WAVELENGTH= 314.90 ANGSTROMS LONG WAVELENGTH= 368.10 ANGSTROMS
X Y Z
SATELLITE POSITION -0.4522744629D 04 -0.3728375571D 04 -0.3344427368D 04 (KMS.)
SATELLITE VELOCITY 0.5448351622D 01 -0.3180418451D 01 -0.1826936021D 01 (KMS./SEC.)
SOLAR POSITION 0.9625782371D 00 -0.2211085070D 00 -0.9588360973D 01 (A.U.)

FIGURE 3.1. COMPUTER OUTPUT OF ORBITAL DATA

$$\frac{d^j x_i}{dt^j} = - \frac{\mu}{R^3} \frac{d^{j-2} x_i}{dt^{j-2}} \quad (3.2)$$

and then employed the truncated Taylor series

$$x_i(t+\Delta t) = x_i(t) + \sum_{j=1}^N \frac{1}{j!} \frac{d^j x_i}{dt^j}(t) \Delta t^j \quad (3.3)$$

The computed elements were employed to evaluate the appropriate altitude and sun angle. In computing the altitude we allowed for the earth's flatness, f , using the relationship

$$R_E = R_0 (1 - f(1 + 1.5f) \sin^2 \phi (1 - \sin^2 \phi / (1 + \frac{2}{3} f^2))) \quad (3.4)$$

where R_E is the distance from the geocenter to the earth's surface at latitude ϕ and R_0 is the equatorial radius. f was taken to be $\frac{1}{298.5}$. R_E changed by about 5 kms. over the orbit.

In Figure 3.2 we present the computed altitude as a function of time for three orders of interpolation (2, 3 and 4). (It should be pointed out that the lack of continuity seen in the $N=2$ data was the worst of any data analyzed to date.) On the basis of these data we have elected to use fourth order near circular orbit interpolation for all future runs.

The sun angle was computed using geocentric solar position coordinates obtained by linear interpolation between the one minute values extracted from the data tape. A plot of

FIGURE 3.2. OSO-7 ALTITUDE VS TIME FOR VARIOUS ORDERS OF INTERPOLATION

ALTITUDE
(KMS)

1972

66TH DAY

N=4

N=3

N=2

ORIGINAL PAGE IS
OF POOR QUALITY

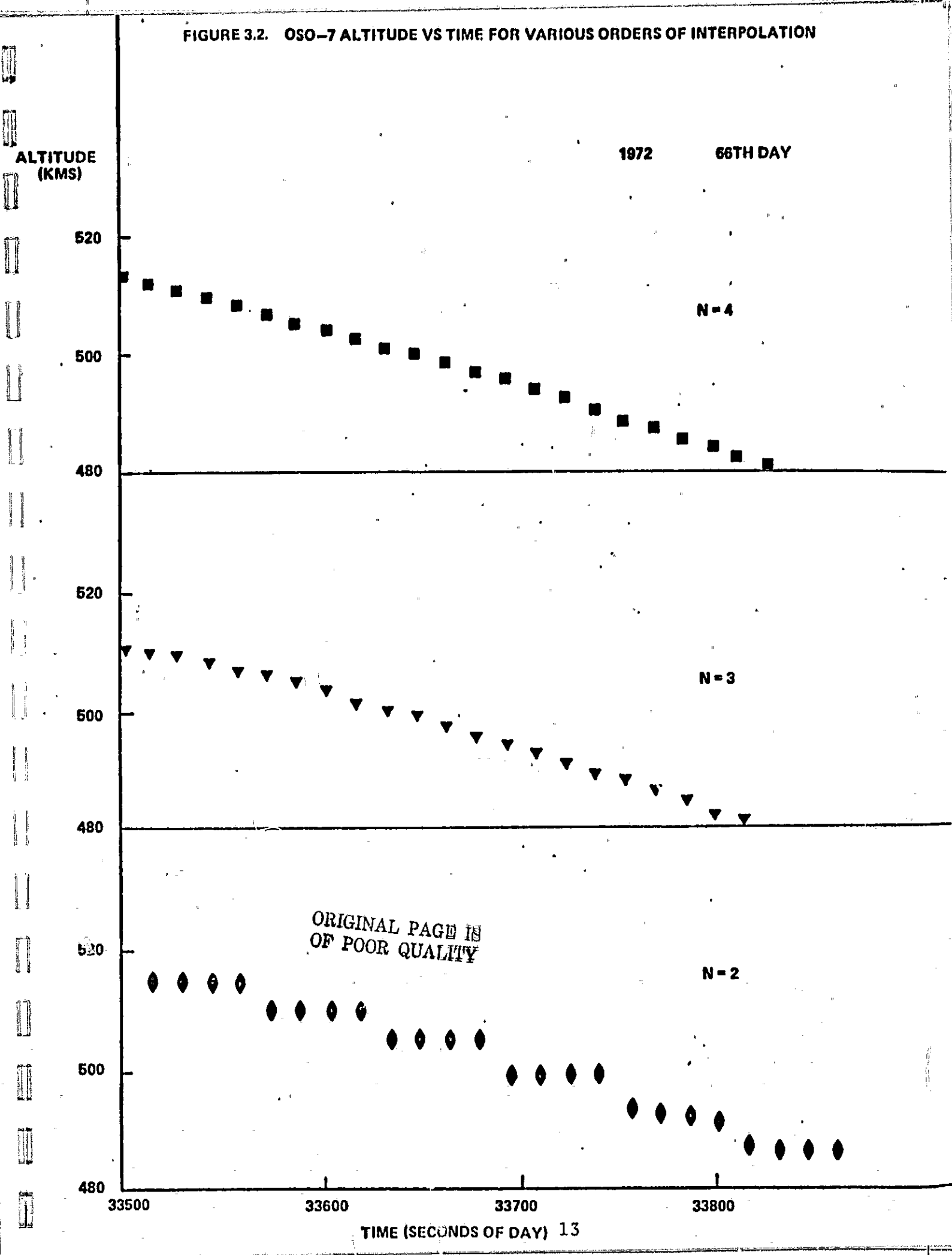
33500

33600

33700

33800

TIME (SECONDS OF DAY) 13



the geocentric solar position coordinates against time indicated that the errors arising from this procedure would be negligible.

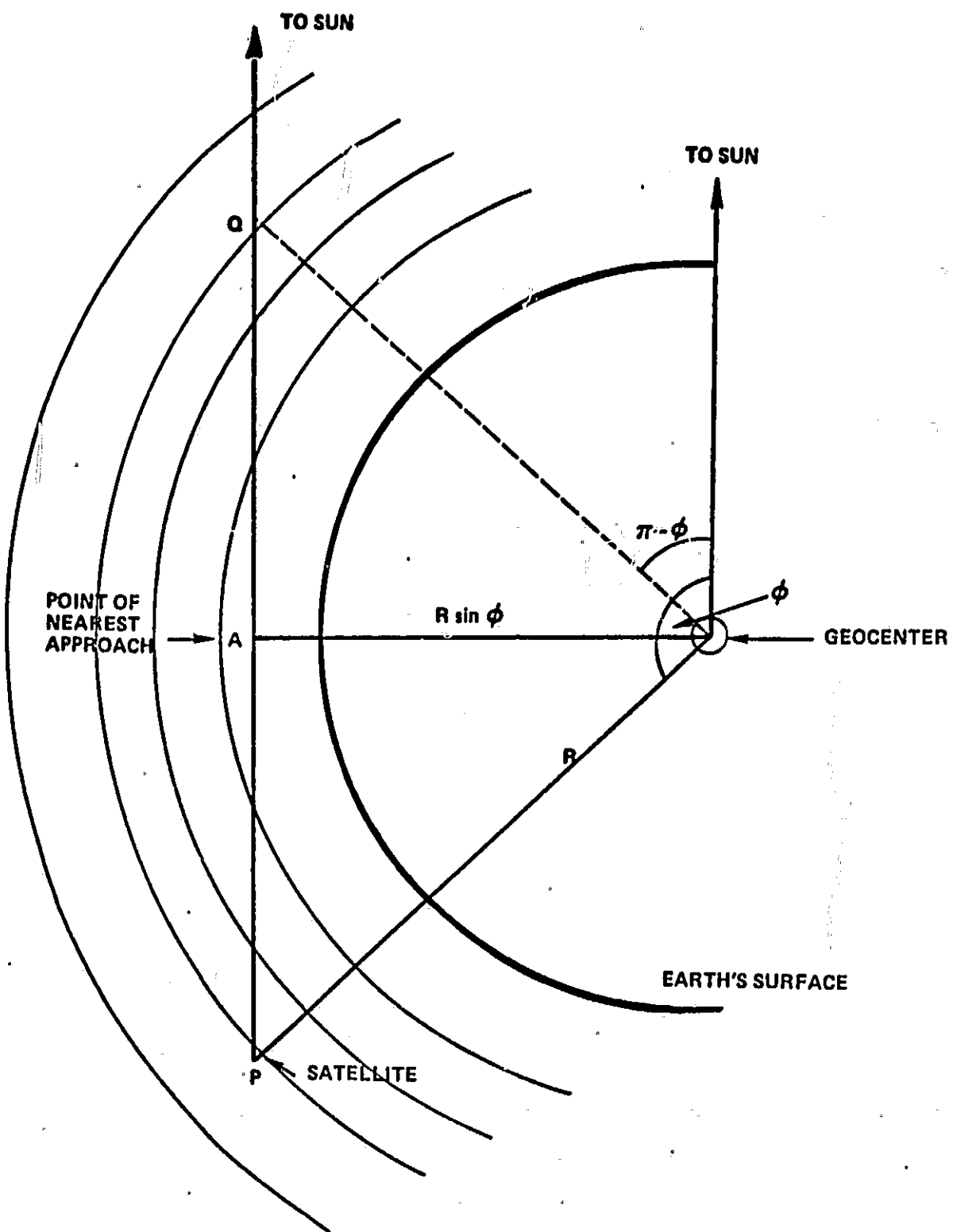
3.2 GEOMETRIC ANALYSIS

In order to analyze the optical thickness data we needed to compute the distance covered through each atmospheric layer of our model. The philosophy employed here was that the earth was spherical, the radius being that appropriate for the latitude of the satellite at the time of data acquisition. This was assumed to give a reasonable representation of the atmospheric layer structure while reducing the computer time for the analysis to a minimum.

Figure 3.3 presents the geometry of the problem in the plane of the geocenter, the satellite and the Sun. P represents the satellite, the solar angle, ϕ , of the satellite being as shown. On the path to the Sun from P, A is the point of closest approach, distant $R \sin \phi$ from the geocenter, where, R is the geocentric radius to the satellite. Using Pythagorus's theorem the distance covered between the Jth and (J+1) the layer beneath the satellite altitude is

$$2 \left[\sqrt{R_{J+1}^2 - R^2 \sin^2 \phi} - \sqrt{R_J^2 - R^2 \sin^2 \phi} \right]$$

whereas the corresponding value for layers above the satellite altitude is the same, but with the 2 replaced by unity. R_J is the geocentric radius of the Jth atmospheric layer.



ORIGINAL PAGE IS
OF POOR QUALITY

FIGURE 3.3. GEOMETRIC ANALYSIS

Clearly, if $R \sin \phi$ is less than $R_E + h_{\min}$ where R_E is the (local) earth radius and h_{\min} the lowest altitude considered and ϕ is greater than $\frac{\pi}{2}$ then no signal can be received, and the data point is ignored.

Suppose that R_K is the radius of the lowest layer boundary crossed by the path. In this case, the distance covered by the path in the Kth layer is just

$$2 \sqrt{R_K^2 - R^2 \sin^2 \phi}$$

If P is in the m th layer then the total distance covered in this layer in the case shown is

$$\sqrt{R_m^2 - R^2 \sin^2 \phi} - 2 \sqrt{R_{m-1}^2 - R^2 \sin^2 \phi} - R \cos \phi$$

The formulae given here were sufficient to evaluate the distance parameters F_{ij} of equation 2.2.

SECTION 4.0

RESULTS OF THE ANALYSIS

The cross-sections per km. were calculated from the data in each layer of the atmosphere from 260 km. to 480 km. The thickness of each layer was 10 km. The total cross-sections from each file for the wavelength of interest are plotted. These are shown in Figures 4.1 to 4.8.

We have fitted the data for 304 and 256 Å models with different exospheric temperatures. The three solid curves shown in the Figures 4.1 to 4.4 are from these models and we see that exospheric temperature 1050° models fits well with the data at sunset at 304 and 256 Å. The total cross-sections obtained from Jacchia models used Knight et al (1973) 256 Å and Hinteregger and Hall (1969) 304 Å absolute cross-sections. We see that the data at lower level layers at sunrise is consistent with sunset data; however, at higher altitudes the sunrise data has large scatter. This scatter possibly can be attributed to some interlocking problems and possibly atmospheric instabilities while heating. Due to lack of knowledge of the cross-sections as a function of temperature and the temperature change itself at the different altitudes, we elected at present to look at the sunset data only.

We derived our estimates of nitrogen and oxygen densities from the best fitted Jacchia model and presented in Table I for 304 and 256 Å. Using these densities in turn at 285 Å and 355 Å, we derived the absolute cross-sections of oxygen and nitrogen at these wavelengths as shown in Section 2.0. The cross-sections thus derived are given in Table II. These cross-sections are new in literature as far as we know. We thus have the densities N_2 and O and cross-sections at 285 Å and 355 Å.

ABSORPTION
CROSS SECTION
(KMS)⁻¹

FIGURE 4.1. TOTAL CROSS SECTION VS. ALTITUDE AT 304 Å⁰
FOR SUNRISE 304 Å⁰

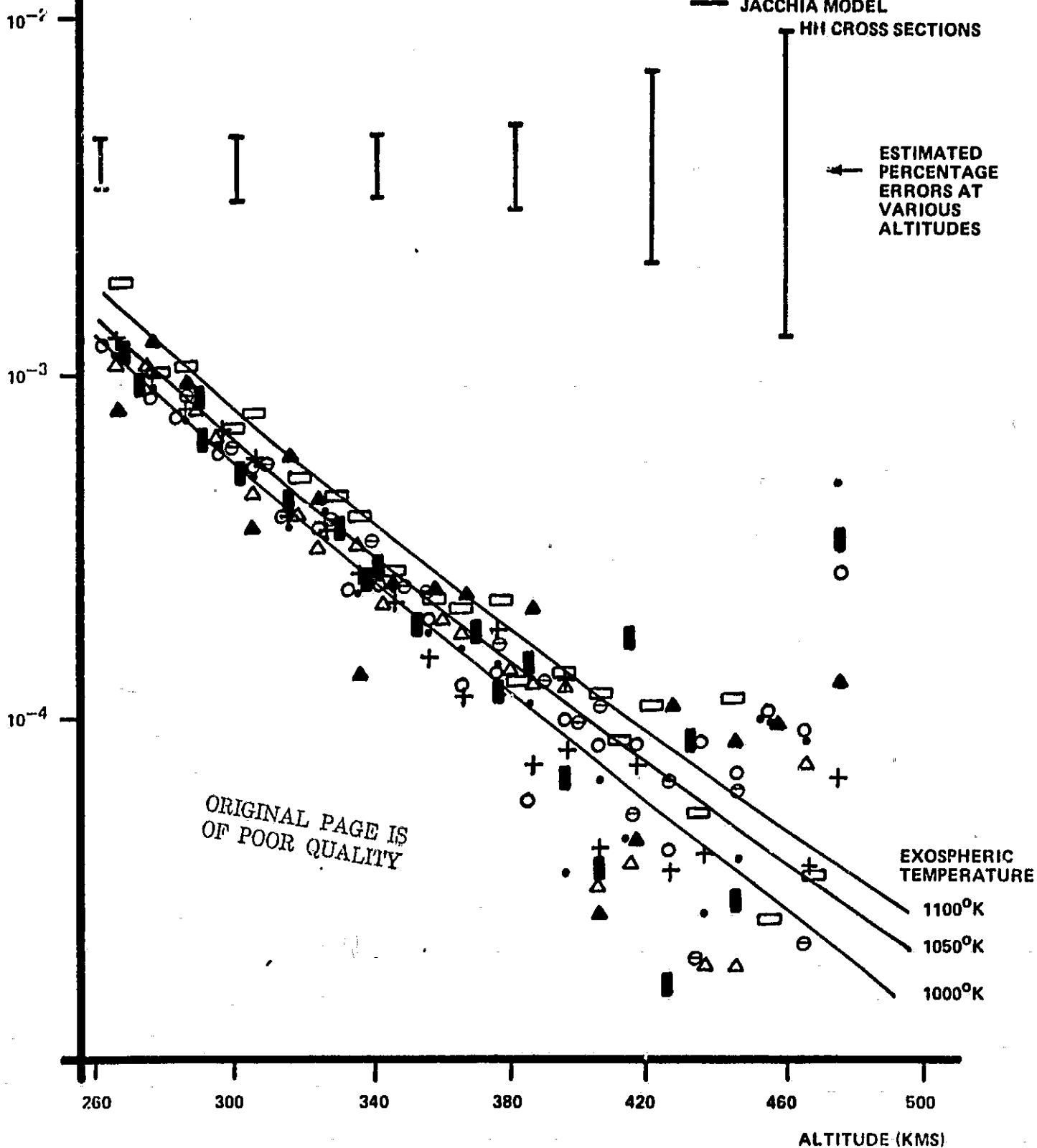
KEY

- CAO631, F18
- + CAO631, F8
- ▲ CAO631, F65
- CAO631, F6
- △ CAO631, F7
- ⊖ CAO629, F15
- CAO631, F27
- CAO630, F15

— JACCHIA MODEL

HH CROSS SECTIONS

ESTIMATED
PERCENTAGE
ERRORS AT
VARIOUS
ALTITUDES



ABSORPTION
CROSS-SECTION
(KMS)⁻¹

FIGURE 4.2. TOTAL CROSS SECTION VS. ALTITUDE AT 256 Å⁰
FOR SUNSET 256 Å⁰

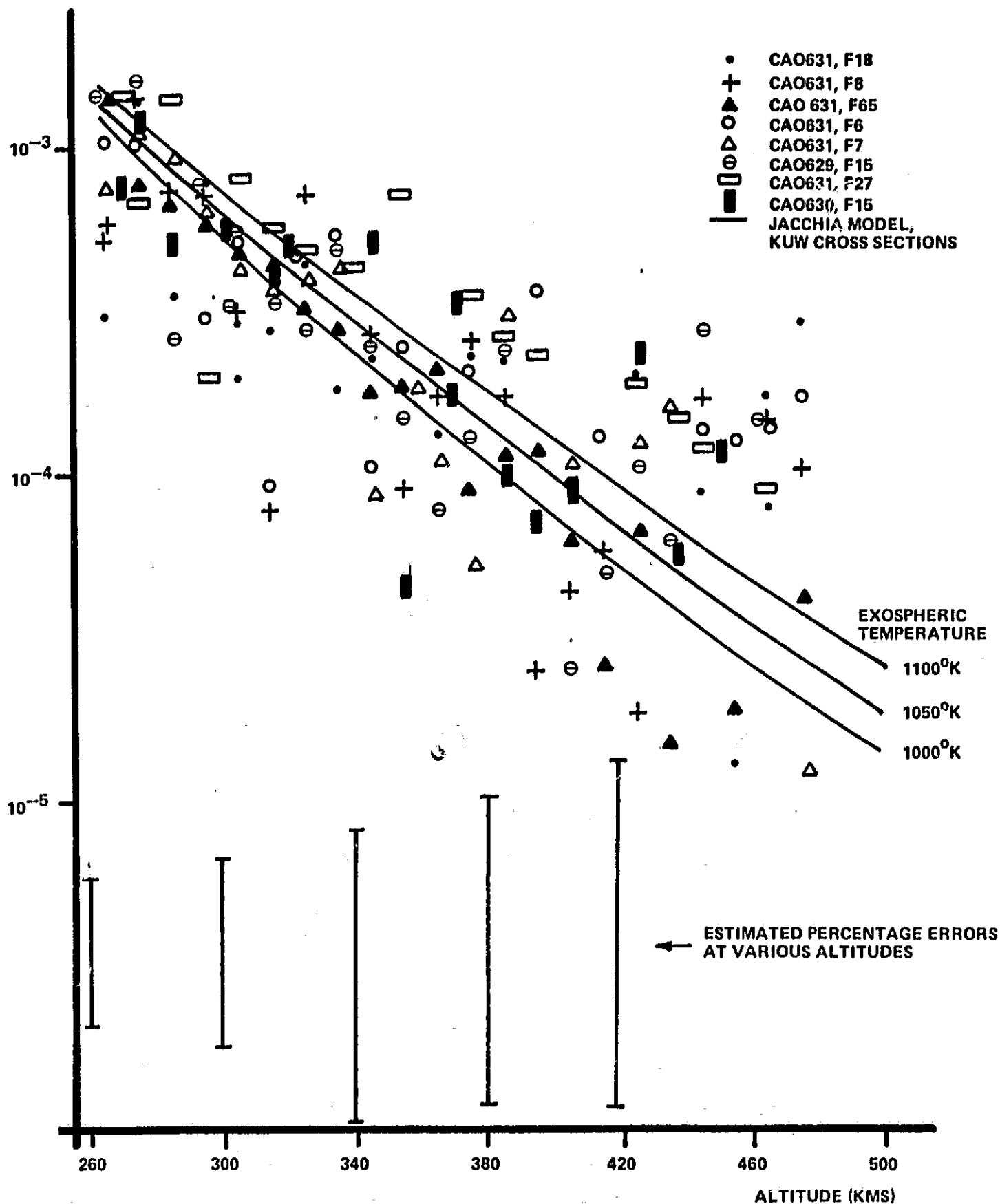


FIGURE 4.3. TOTAL CROSS SECTION VS. ALTITUDE AT 304 Å⁰
FOR SUNRISE 304 Å⁰

ABSORPTION
CROSS SECTION
(KMS)⁻¹

ORIGINAL PAGE IS
OF POOR QUALITY

KEY

- CAO631, F18
- + CAO631, F8
- ▲ CAO631, F65
- CAO631, F6
- △ CAO631, F7
- ⊖ CAO629, F15
- CAO631, F27
- ▬ CAO630, F15
- JACCHIA MODEL,
HH CROSS SECTION

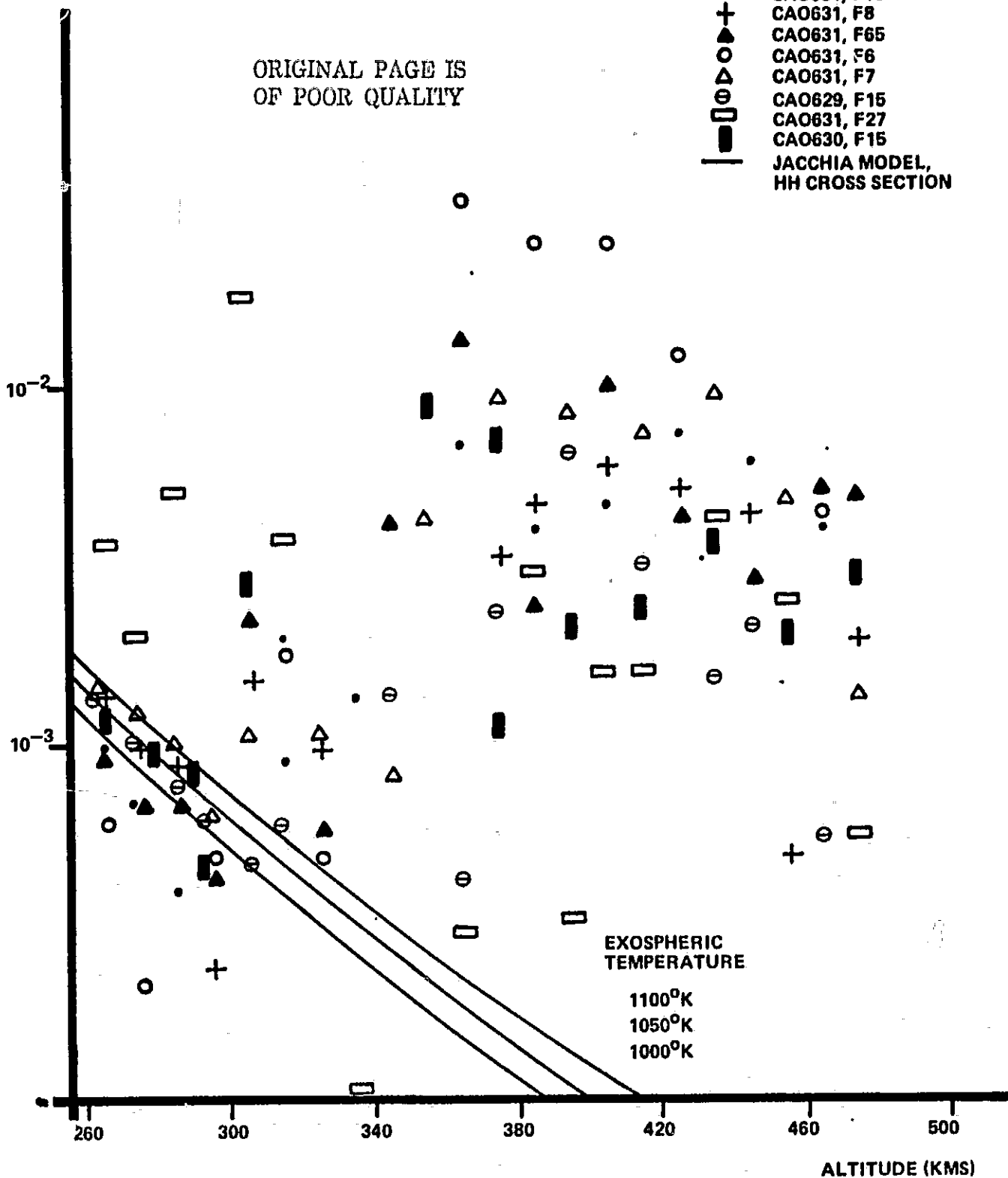


FIGURE 4.4. TOTAL CROSS SECTION VS. ALTITUDE AT 256 Å⁰
FOR SUNRISE 256 Å⁰

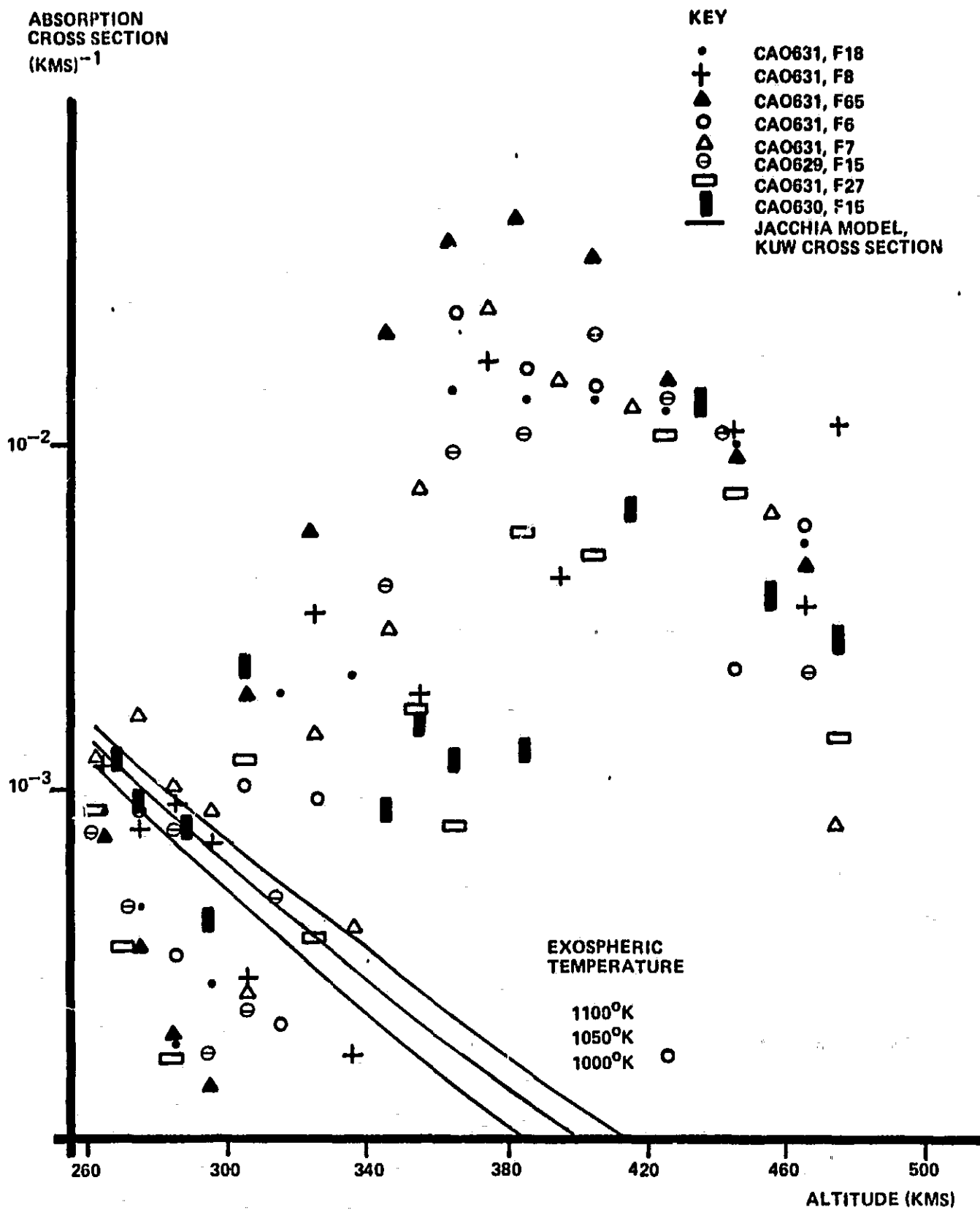


FIGURE 4.5. TOTAL CROSS SECTION VS. ALTITUDE AT 285 Å⁰
FOR SUNSET 285 Å⁰

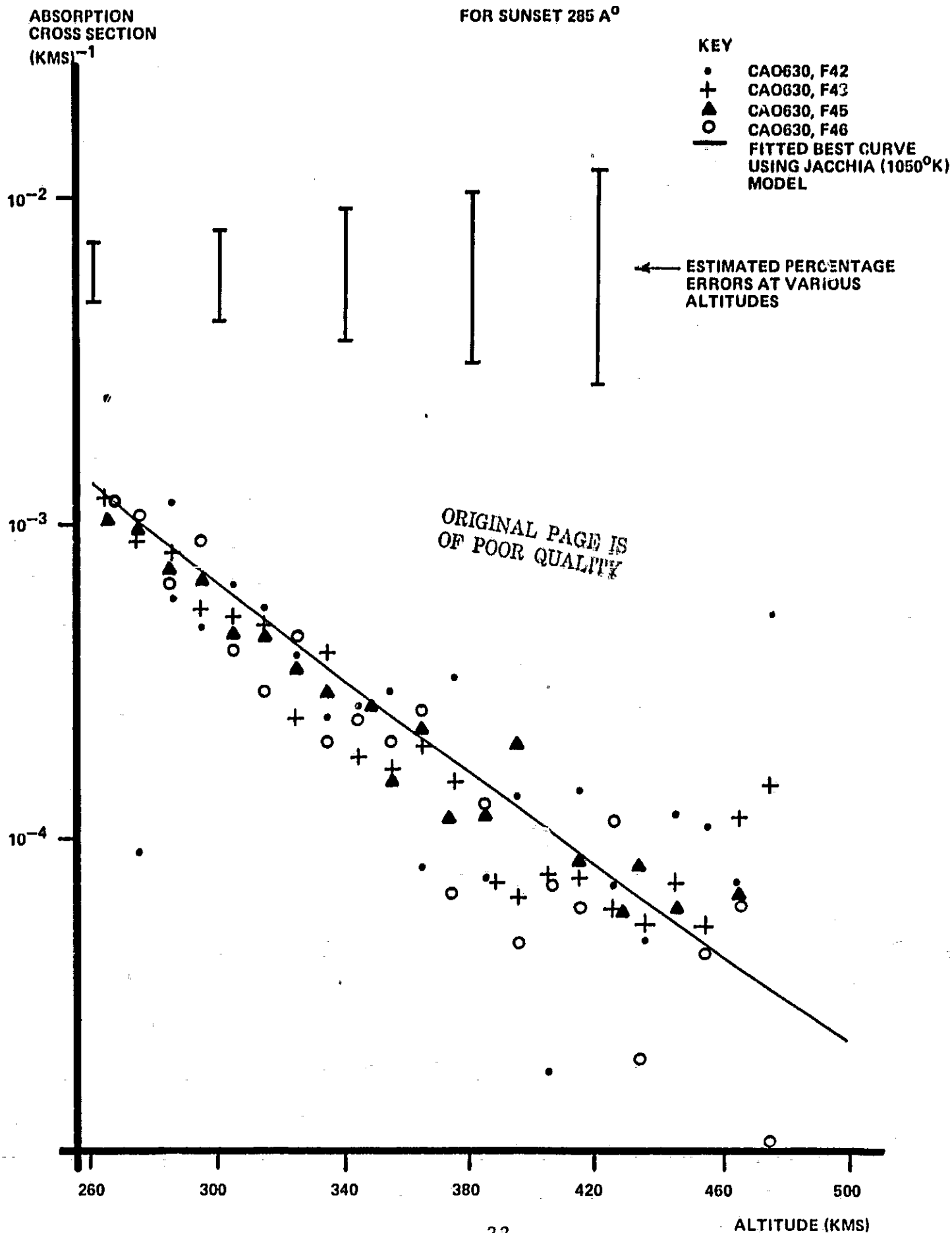
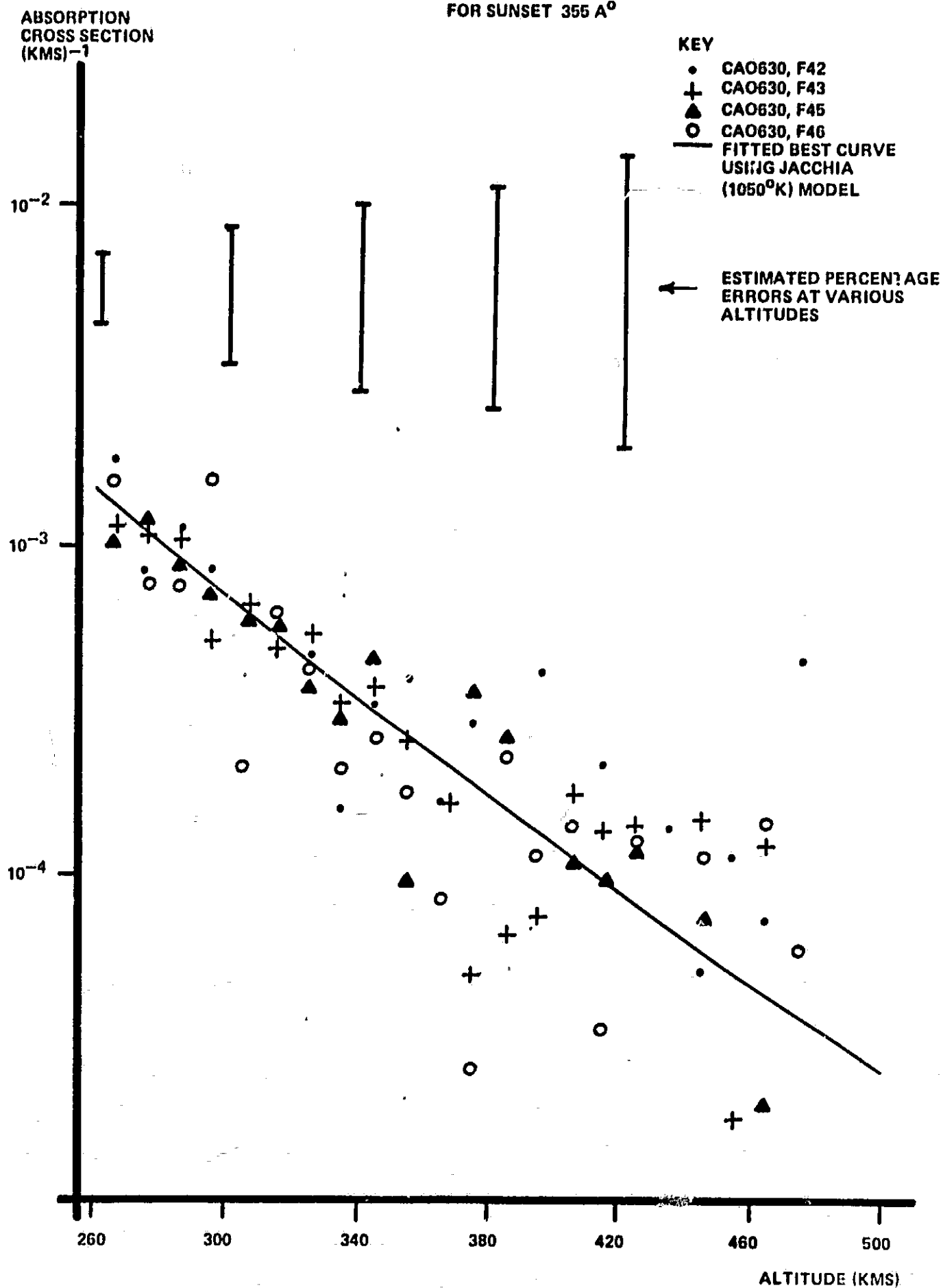
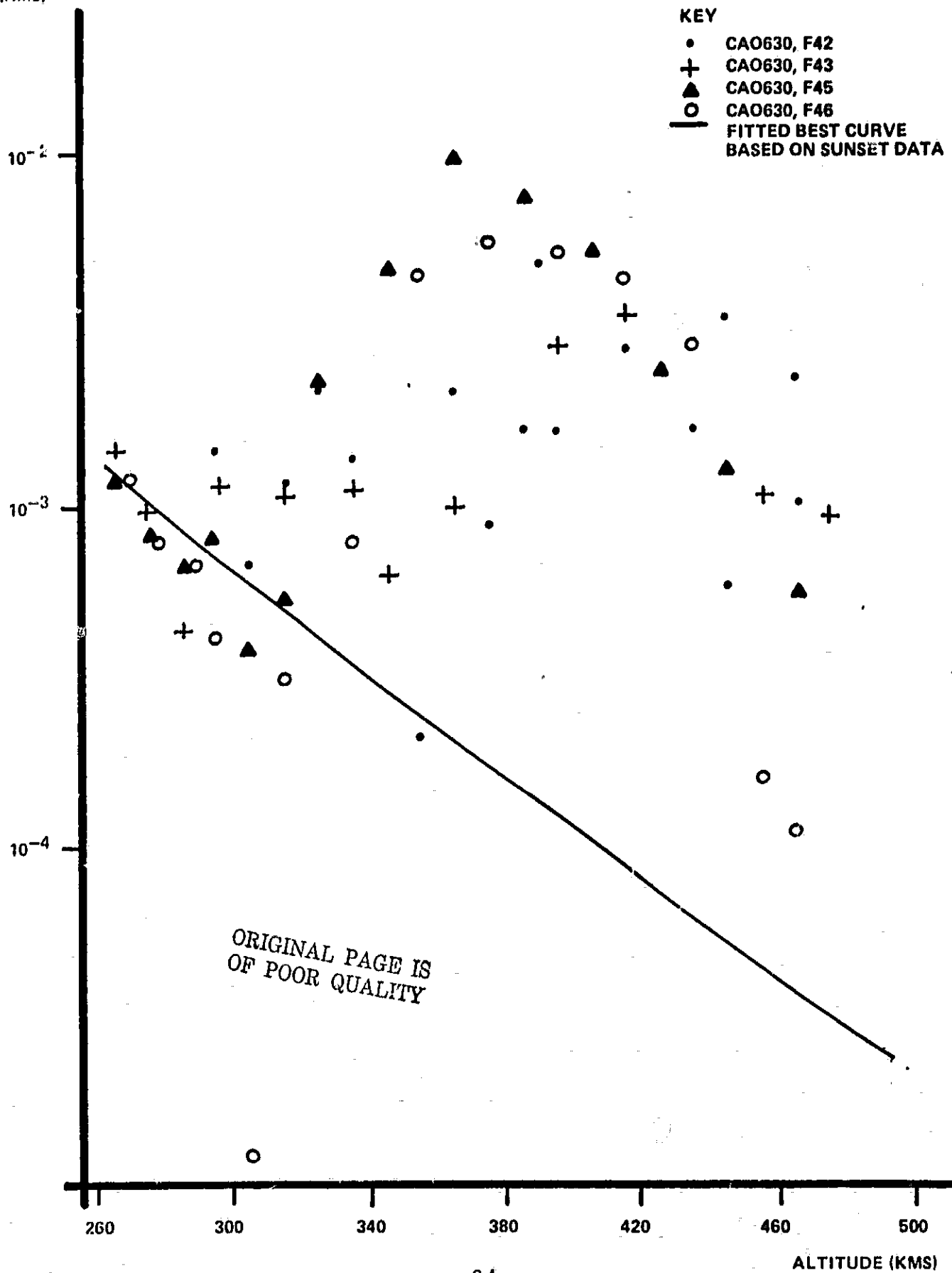


FIGURE 4.6. TOTAL CROSS SECTION VS. ALTITUDE AT 355 Å⁰
FOR SUNSET 355 Å⁰



ABSORPTION
CROSS SECTION
(KMS)⁻¹

FIGURE 4.7. TOTAL CROSS SECTION VS. ALTITUDE AT 285 Å⁰
FOR SUNRISE 285 Å⁰



ABSORPTION
CROSS SECTION
(KMS)⁻¹

FIGURE 4.8. TOTAL CROSS SECTION VS. ALTITUDE AT 355 Å⁰
FOR SUNRISE 355 Å⁰

KEY

- CAO630, F42
- + CAO630, F43
- ▲ CAO630, F45
- CAO630, F46
- FITTED BEST CURVE
BASED ON SUNSET DATA

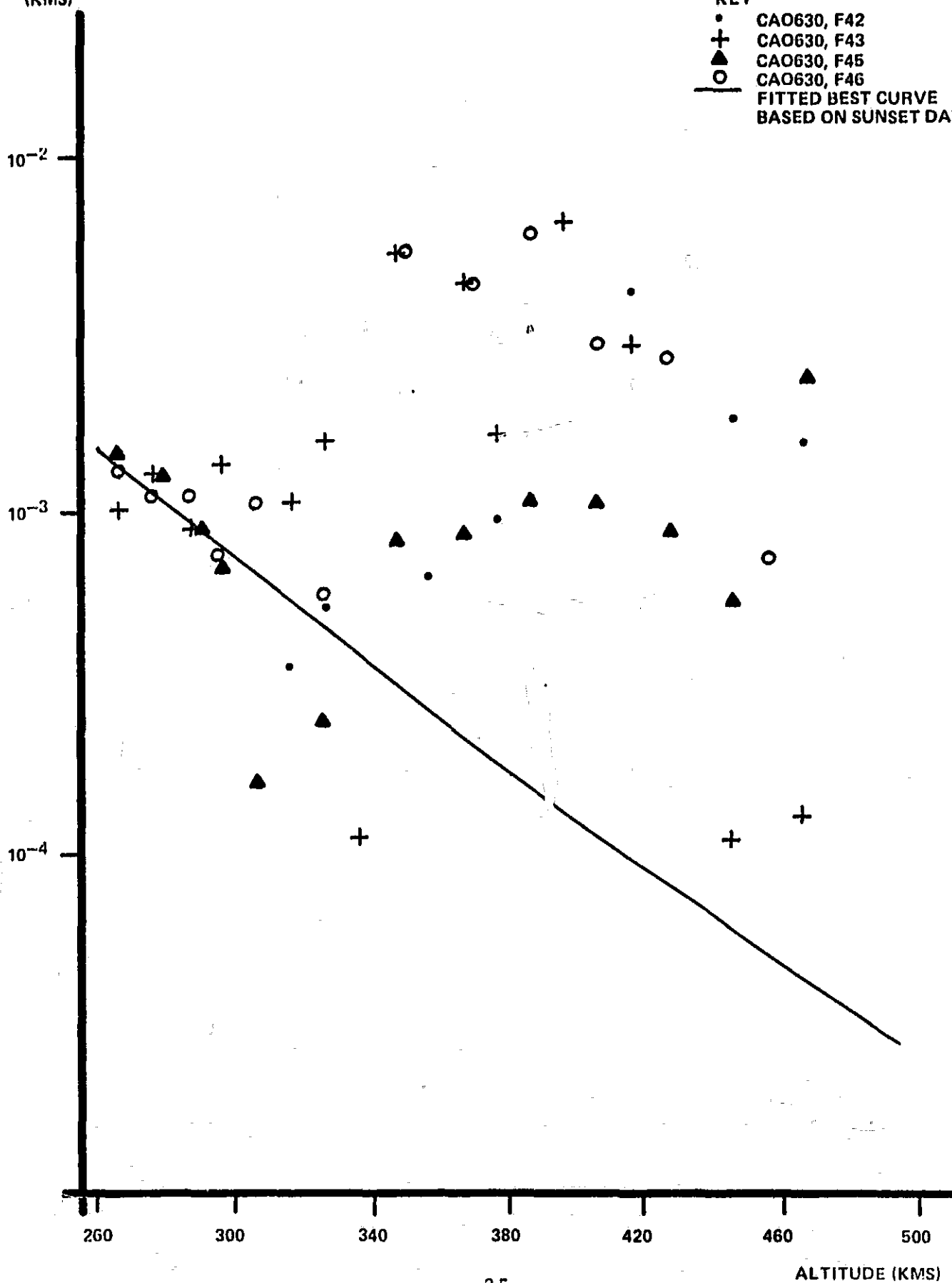


FIGURE 4.9. JACCHIA MODEL AT VARIOUS EXOSPHERIC TEMPERATURES OF O DENSITY VS. ALTITUDE

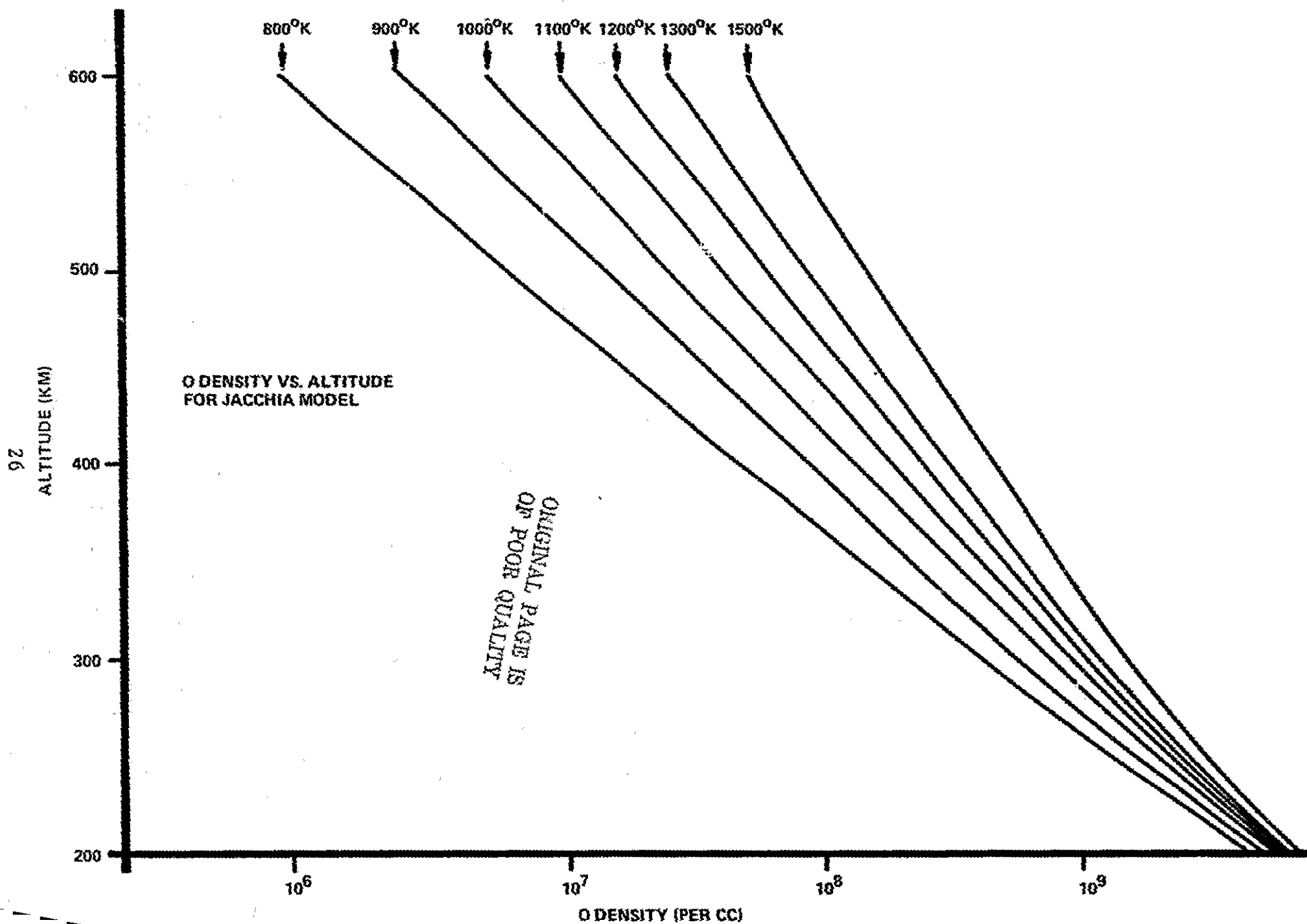
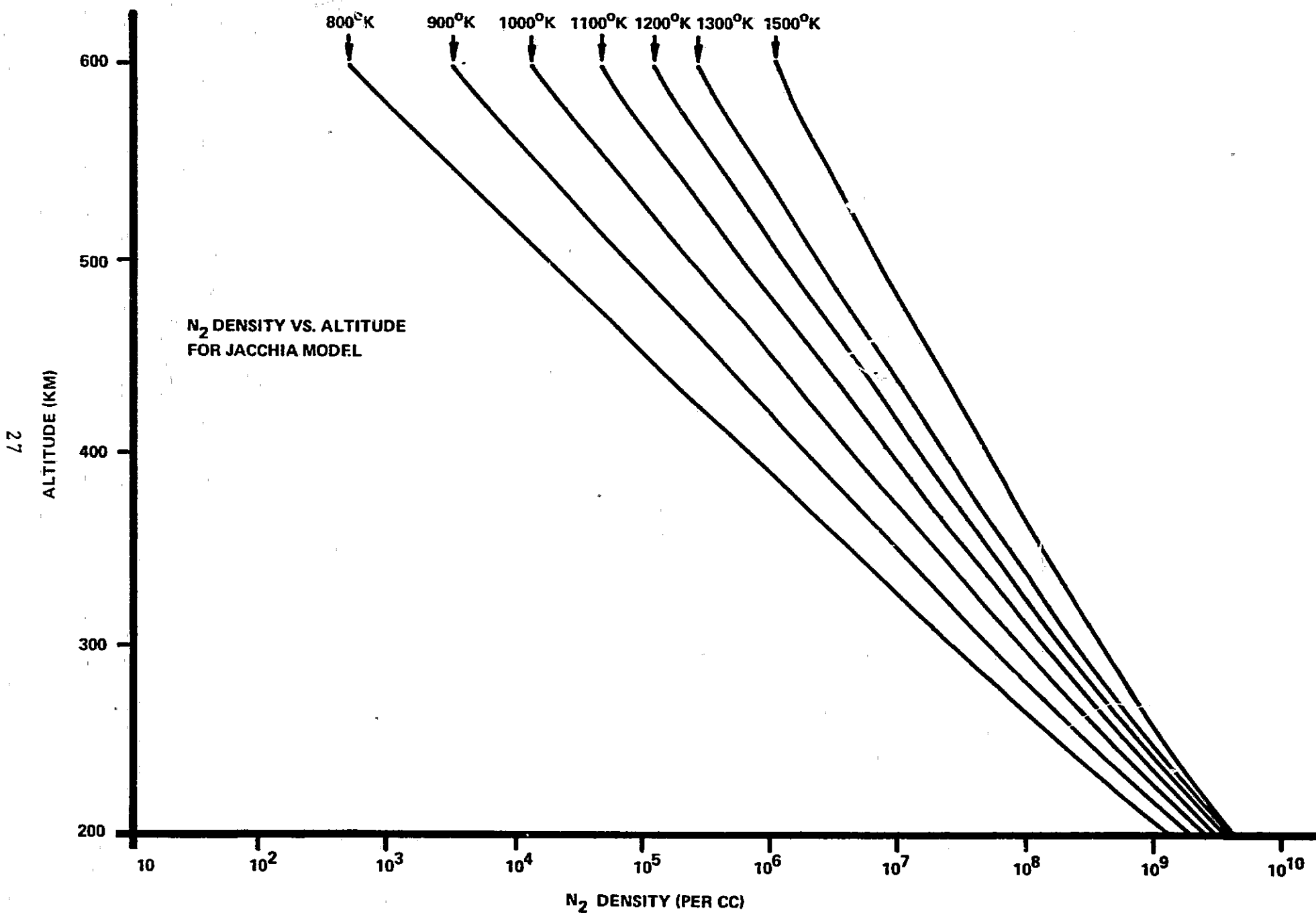


FIGURE 4.10. JACCHIA MODEL FOR VARIOUS EXOSPHERIC TEMPERATURE OF N_2 DENSITY VS. ALTITUDE



In Figures 4.9 and 4.10 we present the oxygen and molecular nitrogen densities versus altitude for different exospheric temperatures of the Jacchia (1971) model that was employed in our study. These graphs illustrate that a significant change in the density and, therefore, the attenuation characteristics of the atmosphere can be expected with changing exospheric temperature. O becomes more predominant with increasing altitude, so information about N_2 must be obtained from attenuation through the lower layers.

SECTION 5.0

DISCUSSION

We have in this project arrived at the goal which is to derive the N_2 , O densities as a function of altitude from 260 km. to 500 km. above the surface of the Earth. However, the data obtained at sunrise is very erratic and showed no trend whatsoever thus giving us only sunset data to deal with. However, the sunset data was adequate to derive the absolute cross-sections at 285 \AA and 355 \AA of O and N_2 which is an unexpected by-product.

There are a number of possible explanations for the erratic data at sunrise. The sunrise data span was much longer in duration than the sunset portion and therefore promises much more information about the atmosphere. It is likely that our description of the atmosphere as varying only in the vertical direction is not valid and that variations in exospheric temperature around the globe become important. There is ample data to justify an extensive survey of this problem but we were unable to study this problem in depth within the time and funding constraints of the project. It is highly desirable to look at attenuation data in this region from other spacecraft. This combining of all the data available from many spacecraft will lead to a high accuracy density profiles and atmospheric model validation.

We thank NASA Headquarters for giving us an opportunity to participate in this program. We, in consultation with the Principal Investigator and the Technical Officer, propose to publish this work in a scientific journal at a later date.

Table I. Densities of Molecular Nitrogen and Atomic Oxygen From the Best Fitted Jacchia Model

NO.	ALTITUDE	DENSITY/CM ³	
		MOL. NITROGEN	OXYGEN
1	260	3.591×10^8	1.676×10^9
2	300	1.048×10^8	8.204×10^8
3	340	3.2×10^7	4.153×10^8
4	380	1.02×10^7	2.146×10^8
5	420	3.28×10^6	1.125×10^8
6	460	1.08×10^6	5.95×10^7
7	500	3.61×10^5	3.18×10^7

Table II. Cross-sections of N_2 and O at
285 Å and 355 Å

CONSTITUENT	CROSS-SECTIONS	
	285 Å	355 Å
MOL. NITROGEN	6.869 \pm 3.0	4.837 \pm 3.9
OXYGEN	6.282 \pm 0.6	7.612 \pm 0.8

UNIT = Megabarn = 10^{-18}cm^2

REFERENCES

- 1969 Hinteregger, H.E., and Hall, L.A., "Thermospheric Densities and Temperatures from EUV Absorption Measurements by OSO III," Space Research, 9, 519-529.
- 1971 Jacchia, L.G., "Revised Static Models of the Thermosphere and Exosphere With Empirical Temperature Profiles," S.A.O. Special Report 332.
- 1973 Knight, D.E., Uribe, R. and Woodgate, B.E., "Low Latitude Density Variations in the Atmosphere Between 200 and 400 km, from August 1969 to May 1970," Planet. Space. Sci., 21, 253-271.